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Anexo A

Trabalhos Publicados

Neste anexo são incluídos os trabalhos publicados pelo autor e seus orientadores ao longo do desenvolvimento da tese:

a) *“Dynamic Characteristics of Rainfall Rate in the Amazon Region and their Effects on Satellite Link Design”*, apresentado no *International Microwave and Optoelectronics Conference*, realizada em Brasília, Brasil, Julho de 2005;

b) *“Precipitation Studies for Rain Attenuation Prediction in The Amazon Region”*, apresentado no *International Symposium on Antennas and Propagation – ISAP 2005*, realizado em Seul, Coréia do Sul, Agosto de 2005;

c) *“Spatial Variability of Rain in the Amazon Region with Application to Site Diversity in Earth-satellite Paths”*, apresentado no *ClimDiff 2005*, realizado em Cleveland, Estados Unidos, Setembro de 2005.

d) *“An Exceptional Rainfall Event in the Amazon Region”*, apresentado no *ClimDiff 2005*, realizado em Cleveland, Estados Unidos, Setembro de 2005.

e) *“Análise da Variação Espacial da Taxa de Precipitação com Aplicação em Diversidade de Sítio para Enlaces Terra-Satélite”*, apresentado no XXII Simpósio Brasileiro de Telecomunicações – SBrT 2005, realizado em Campinas, Brasil, Setembro de 2005;

f) *“Rain Attenuation Research in Brazil”*, apresentado na *XXVIIIth URSI General Assembly – URSI 2005*, realizada em Nova Delhi, Índia, Outubro de 2005;

Dynamic Characteristics of Rainfall Rate in the Amazon Region and their Effects on Satellite Link Design

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Abstract — This paper deals with an investigation of dynamic characteristics of rainfall rate in the Amazon Region. Experimental data from 3 (three) selected locations pertaining to different climate subtypes identified in this tropical rainy region constitute the basis of the analysis carried out here. Details about the annual clock hourly rainfall accumulation and the average duration of rainfall events are presented, the relevance of their effects in satellite system design being pointed out.

Index Terms — Rain, precipitation rate, dynamics properties of rain, worst-month distribution, Köppen climatic classification, satellite communication systems.

I. INTRODUCTION

The evaluation of fade margin due to rain is normally based on the annual cumulative distribution of precipitation rate corresponding to the area under study. This distribution is adequate for planning high availability systems, where large levels of fade margin are used. The same is not true for low availability links as broadcast satellite services (DTH – Direct To Home) or VSAT (Very Small Aperture Terminal) systems. In this situation, where an acceptable outage is around 0.3% of year, the time behaviour has an important impact on system design. Consequently, more detailed information on the dynamics of rain phenomenon is needed, the knowledge of seasonal and diurnal variation of fade duration statistics and fade interval characteristics being of fundamental relevance.

It is recognized that there is no accepted model for the conversion of rain rate duration to slant path duration statistics [1]. The main problem is the possibility of having more than one rain cell along the propagation path. However, according to Timothy et al. [2], for elevation angles higher than 30°, the measured rainfall rate threshold

at a given probability level may be used to obtain the associated slant path attenuation. A study by Lekla and Prapinmongkolkarn [3] corroborated the above conclusion. Of course, the accuracy of this procedure depends on the model used for the evaluation of rain attenuation. In this context, this paper discusses seasonal and diurnal measurements of rainfall rate carried out in the Amazon region. Rainfall data presented here correspond to preliminary results from a research program on rain attenuation which is being developed under the responsibility of the Military Institute of Engineering with financial support from the National Scientific and Technological Development Council (CNPq).

II. CLIMATES IN THE AMAZON REGION

According to Köppen [4], the climate of Amazon region is a tropical rainy type (A) where 3 subtypes can be identified:

- a) Rainy equatorial (Af) – with a large annual rainfall (over 2000 mm) and practically no dry season;
- b) Monsoon tropical (Am) – the annual rainfall is equal to or larger than Af, but there is a short dry season (one to three months);
- c) Wet-and-dry tropical (Aw) – where the rain and dry seasons are well defined.

Fig. 1 shows the geographical limits of these climatic subtypes and the rain gauge network implemented in the Amazon region. Tipping bucket gauges with 0.1 mm capacity and 1 minute integration time are being used. Typical average monthly rainfall amount (mm) from 3 selected locations (Cruzeiro do Sul – Af, Santarém – Am

and Boa Vista – Aw) are depicted in Fig 2. It is clear from this figure that the subtype Am is a transition climate between Af and Aw.

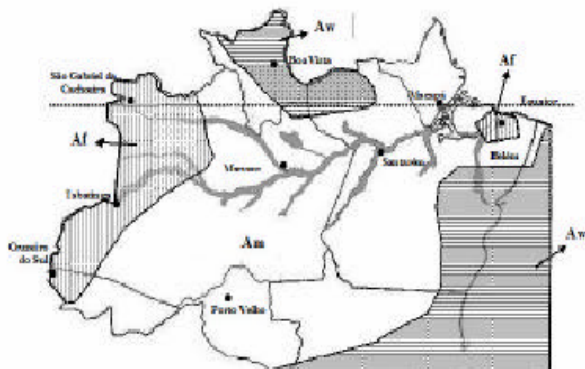
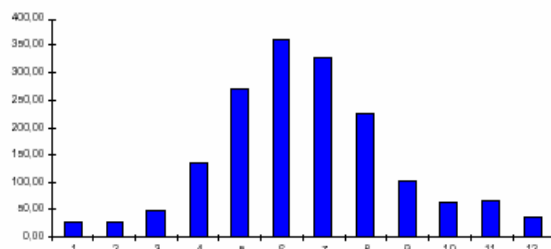
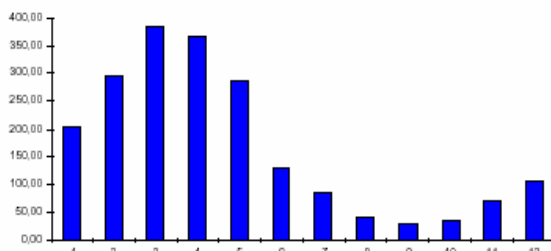


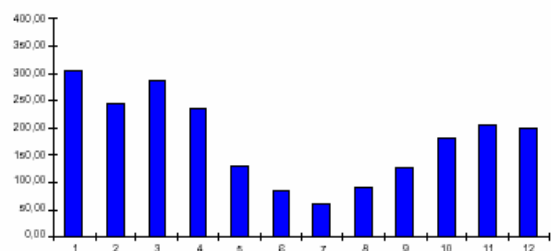
Fig. 1. Köppen climate classification in the Amazon region.



(a)



(b)

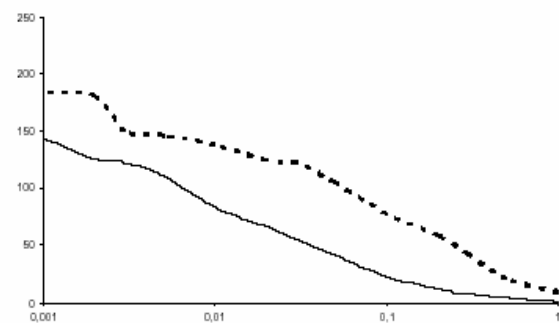


(c)

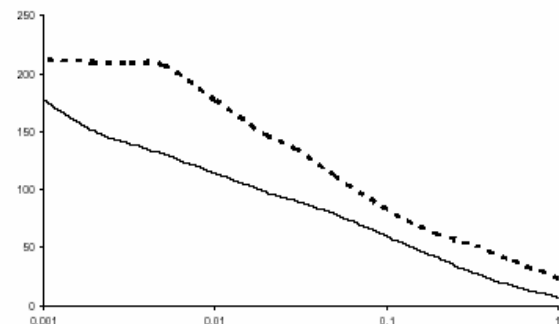
Fig. 2 Typical average monthly rainfall amount (mm) (a) Boa Vista (S 2.79 W 54.72); (b) Santarém (S 2.50 W 54.72); (c) Cruzeiro do Sul (S 7.61 W 72.68).

III. DYNAMIC BEHAVIOUR OF RAINFALL RATE

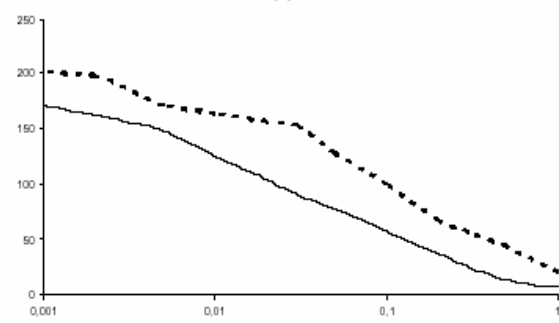
Besides the average annual outage due to fading, another criterion used in the planning of radio communication links is the average of worst month [5] statistics (see Fig. 3). According to this criterion, typical outages of low availability systems are around 1 to 0.1% of the worst month. It is observed in Fig. 3 that this percentage range corresponds to rainfall rates between 20 and 100 mm/h. In spite of being more rigorous with this procedure, as it will be shown below, the analysis of seasonal and diurnal variations should be taken also into account.



(a)



(b)



(c)

Fig. 3 Annual (—) and worst-month (---) rainfall rate distributions (mm/h). (a) Boa Vista; (b) Santarém; (c) Cruzeiro do Sul.

Fig. 4 shows the annual clock hourly rainfall accumulation for Cruzeiro do Sul, Santarém and Boa Vista. It is observed that the concentration of rain varies along the day from one location to another. Particularly important is the maximum accumulation in Cruzeiro do Sul, which occurs in the afternoon and early evening and Santarém from late evening to morning hours. Of course, depending on the service being planned, this behaviour may be quite important. For instance, intense rain in the morning or afternoon (business hours) affects commercial services, while the evening is the prime time for radio broadcasting. It should be pointed out that the precipitation accumulation along the day is directly connected to the clock hourly rainfall rate exceedance. An example is given in Fig. 5.

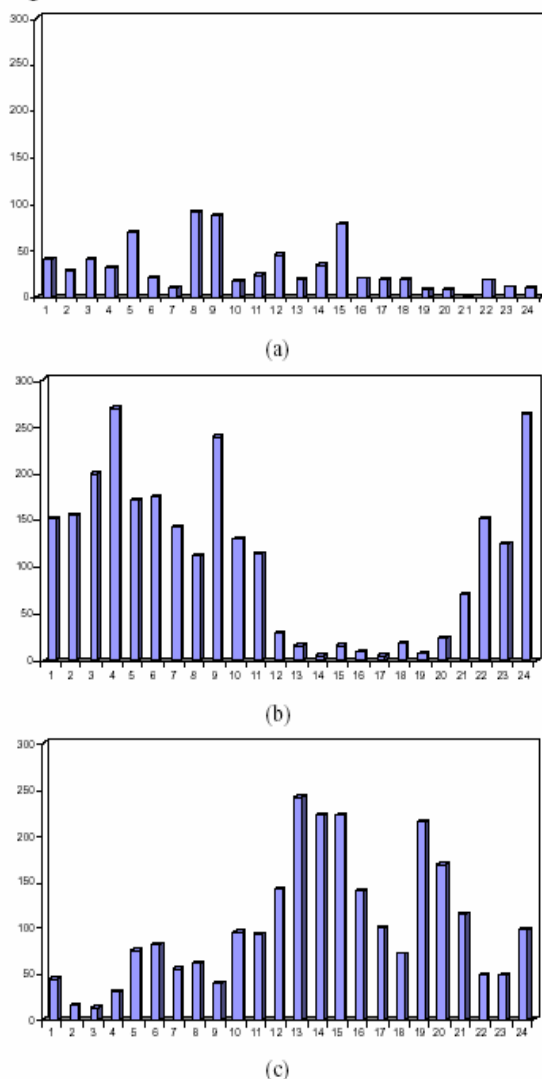


Fig. 4 Annual clock hourly rainfall accumulation (mm). (a) Boa Vista; (b) Santarém; (c) Cruzeiro do Sul.

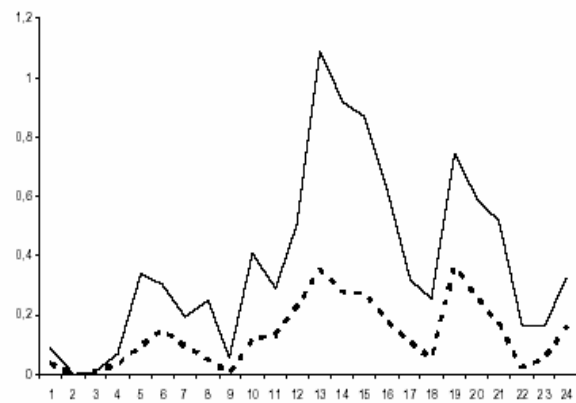
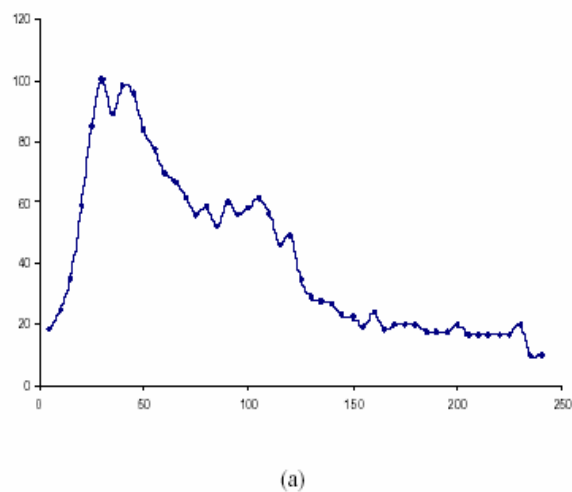
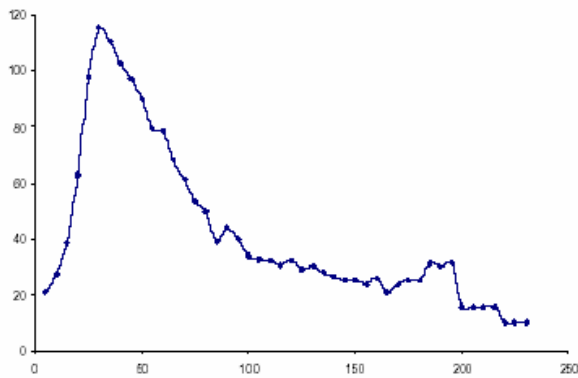


Fig. 5. Annual clock hourly exceedance (%) – Cruzeiro do Sul – 20 mm/h (—) and 50 mm/h (---).

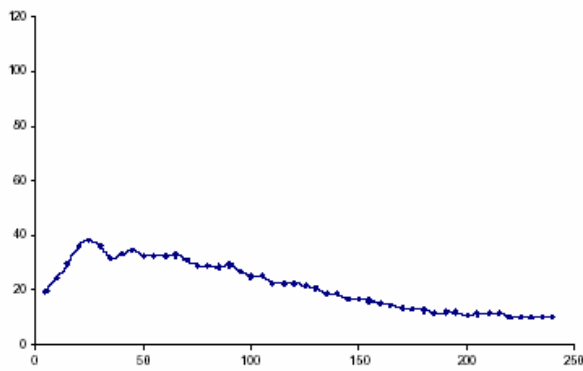
Fig. 6 depicts the average duration of rainfall events as a function of precipitation rate for the locations under study. In all locations, long events are concentrated from 20 to 70 mm/h. Incidentally, these values are within the range from 1 to 0.1% of the worst month statistics commented before (see Fig. 3). The design of mitigation techniques, as forward error correction codes (FEC), uplink power control, etc., normally used to overcome fading effects, should take these data into consideration. Other dynamic characteristics such as seasonal variations (wet and dry seasons) and the average interval between rainfall events, not shown here due to paper length limitation, are also important for satellite system planning.



(a)



(b)



(c)

Fig. 6. Average duration of rainfall events (s) as a function of precipitation rate (mm/h) (a) Boa Vista; (b) Santarém; (c) Cruzeiro do Sul.

IV. CONCLUDING REMARKS

Dynamic characteristics of rainfall rate in the Amazon region were investigated in this paper. The behaviour of precipitation in three locations pertaining to different climate subtypes identified in this tropical rainy region was described and some relevant aspects of interest for satellite system design were pointed out. The analysis carried out here was based on the hypothesis that, for high elevation angles, it is possible to convert rain rate duration to slant path duration statistics. Fortunately, this is the case of greatest interest in the Amazon region. Anyway, the research is progressing and more data will be published shortly, including comparison with rain attenuation measurements.

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PRECIPITATION STUDIES FOR RAIN ATTENUATION PREDICTION IN THE AMAZON REGION

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1. Introduction

Rainfall rate is an essential input parameter in the prediction of rain attenuation. Mathematical models used in this prediction are function of precipitation rate distribution at a point in or near to the propagation path. However, such distributions are available for a limited number of locations. A solution to overcome this difficulty is the extrapolation from one point to another one belonging to the same climatic region. In this context, rainfall characteristics of the Amazon region is described in this paper. Rainfall data presented here correspond to preliminary results from a research program on rain attenuation which is being developed under the responsibility of the Military Institute of Engineering with financial support from the National Scientific and Technological Development Council (CNPq).

2. Köppen Climate Classification

The structure of Köppen climatic classification depends on temperature, precipitation and vegetation [1]. Once these meteorological factors can be related to the statistical distribution of rainfall rate, Köppen classification was adopted in this paper. According to this classification, the climate of Amazon region is a tropical rainy type (A) where 3 subtypes can be identified:

- a) Rainy equatorial (Af) – with a large annual rainfall (over 2000 mm) and practically no dry season;
- b) Monsoon tropical (Am) – the annual rainfall is equal to or larger than Af, but there is a short dry season (one to three months);
- c) Wet-and-dry tropical (Aw) – where the rain and dry seasons are well defined.

Figure 1 shows the geographical limits of these climatic subtypes and the rain gauge network implemented in the Amazon region. Tipping bucket gauges with 0.1 mm capacity and 1 minute integration time are being used. Typical average monthly rainfall amount (mm) and temperature ($^{\circ}\text{C}$) from 3 selected locations (Boa Vista – Aw, Santarém – Am and Cruzeiro do Sul – Af), are shown in Figure 2 [2]. It is clear from this figure that the subtype Am is a transition climate between Af and Aw. However, as it will be commented in the next Section, due to higher rainfall rates in the rainy season, its effect on system availability is quite important.

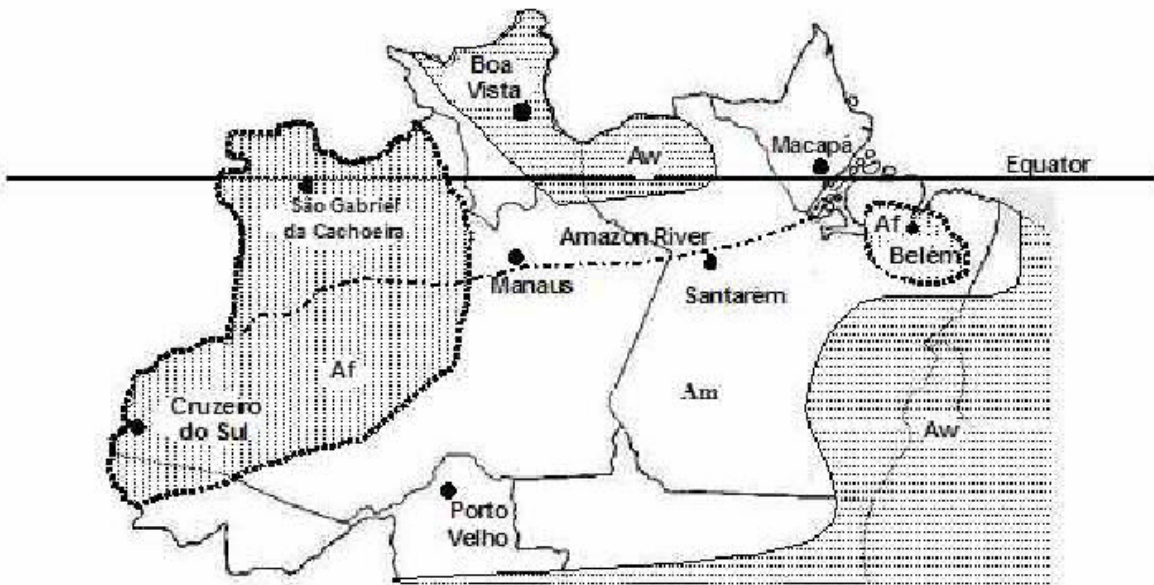
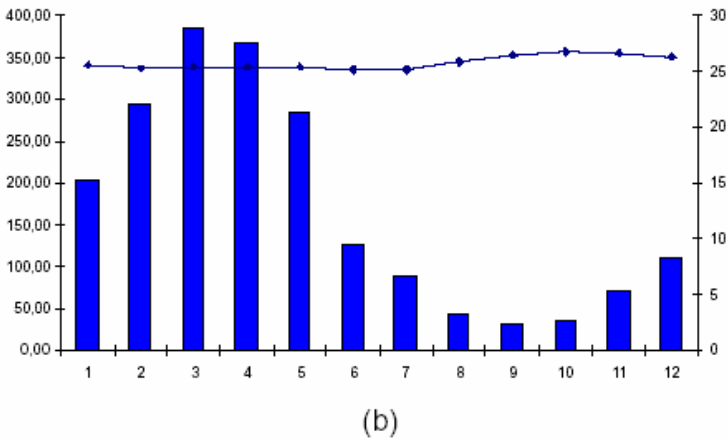
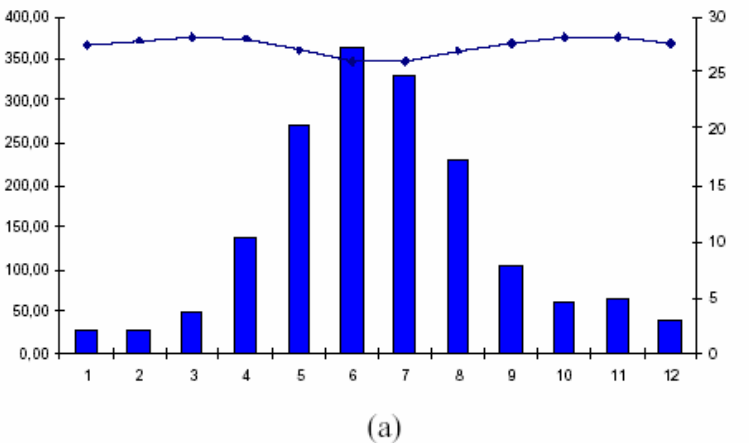


Figure 1 – Köppen climate classification in the Amazon region.



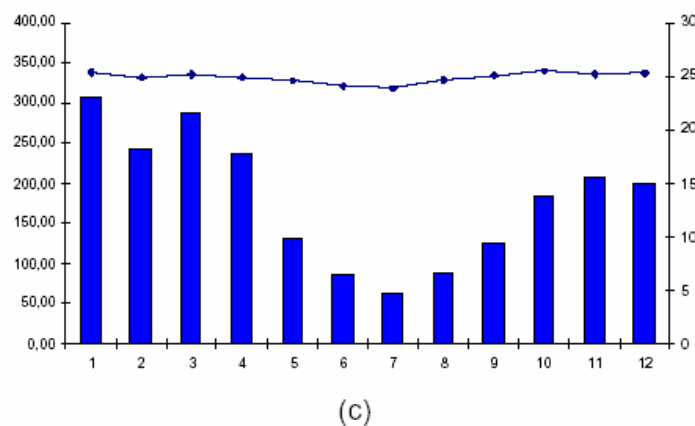


Figure 2 – Typical average monthly rainfall amount (mm) and temperature ($^{\circ}\text{C}$) in the Amazon region. (a) Aw - Boa Vista (S 2.79 W 54.72); (b) Am – Santarém (S 2.50 W 54.72); (c) Af – Cruzeiro do Sul (S 7.61 W 72.68).

3. Rainfall Characteristics in the Amazon Region

Figures 3 and 4 show the annual and the monthly cumulative distributions for the above cited locations. As expected, based on the total annual amount of rain (over 2000 mm for both Cruzeiro do Sul and Santarém), there is no significant difference in the annual cumulative distribution of rainfall rate between Af and Am subtypes. On the other hand, in the case of Boa Vista (subtype Aw) lower rain rates are observed.

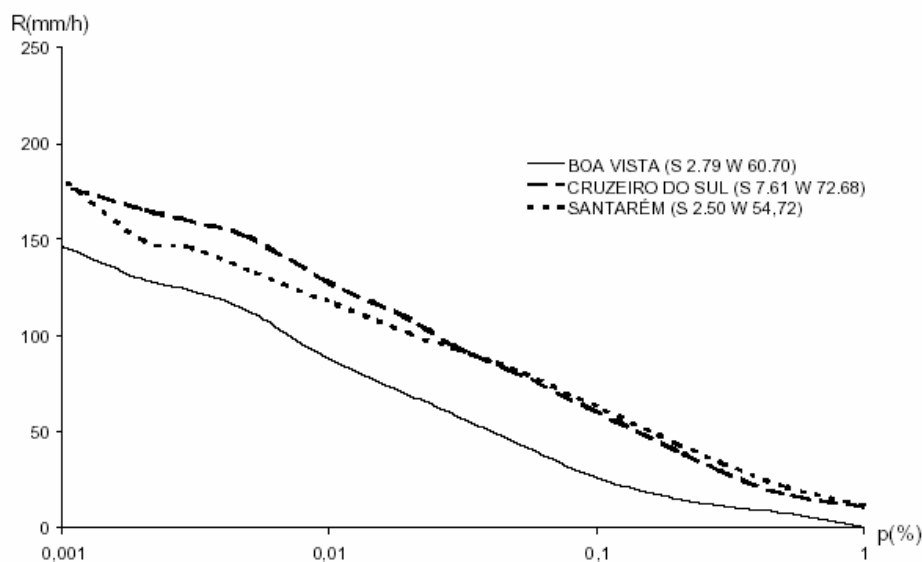


Figure 3 – Annual rainfall rate distributions.

The statistical behaviour is not the same when considering month distributions [3]. For relatively low rainfall rates, i.e. in the range from 1 to 0.1 percent of time, distributions for Af and Am subtypes are almost the same. This behaviour is of interest for low availability systems as, for instance, satellite broadcasting reception. However, from 0.01

to 0.001 percent of time, higher rainfall rates are observed in the subtype Am. In this subtype there is a large difference between dry and rainy seasons. When planning high availability communication systems in this climate, this behaviour should be taken into account. A similar conclusion appears in the case of subtype Aw. Possibly, automatic transmission power control (ATPC) can be used as mitigation technique to compensate for such high rainfall rates.

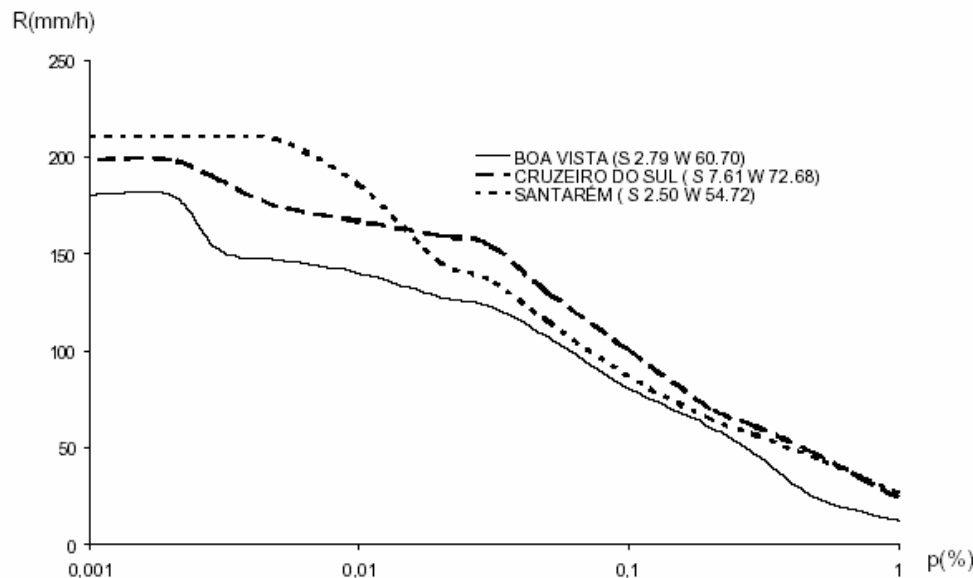


Figure 4 – Worst-month rainfall rate distributions.

4. Concluding Remarks

Based on Köppen climate classification, rainfall characteristics of Amazon region were described in this paper. In spite of being classified as a tropical rainy climate, 3 subtypes were identified (Af, Am and Aw). Under the practical point of view, the importance of each subtype depends on the particular aspect being considered. Although, regarding the annual cumulative distribution, each one follows the expected behaviour, the same is not true for the worst month distribution. In this context, different procedures should be taken into account when planning low or high availability communication systems.

5. References

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Spatial Variability of Rain in the Amazon Region with Application to Site Diversity in Earth-Satellite Paths

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Abstract

This paper deals with the spatial variability of rain rate in the Amazon region. Measurements from a meteorological radar located in Cruzeiro do Sul [7.36°S ; 72.46°W] were taken into account and a mathematical model for relating the statistical distribution of precipitation rate between two points separated by a given distance was developed. Data from a pluviograph installed near the radar was used in the derivation of a relationship between radar reflectivity and precipitation rate. Final results were promising enough. The behavior of the mathematical model has showed excellent agreement with site diversity gain curves from empirical models based on rain attenuation.

1. Introduction

This paper investigates the spatial variability of rain rate in the Amazon region aiming to derive a physical model to evaluate the performance of site diversity in earth-satellite paths. In this context, measurements from a meteorological radar located in Cruzeiro do Sul [7.36°S ; 72.46°W], in the state of Acre, Brazil, were taken into account and a mathematical model relating the statistical distribution of precipitation rate between two points separated by a given distance was developed. Data from a pluviograph installed near the radar was used in the derivation of a relationship between radar reflectivity and precipitation rate. Experimental data presented here are from a research program which is being developed under the responsibility of the Military Institute of Engineering (IME) with financial support from the National Scientific and Technological Development Council (CNPq). Figure 1 shows the rain gauge and radar stations used in this program, which are co-located and covering all the Amazon region.



Figure 1. Amazon region – rain gauge and radar stations.

2. Mathematical model

The fundamental problem to be solved by this mathematical model is the evaluation of the separation distance L between two earth stations, in a site diversity configuration, aiming to minimize the rain effect and in accordance with a specified system availability criterion. This task can be performed having as statistical basis the probability of a rainfall rate R_0 be exceeded, simultaneously, in two points separated by L . This probability $P(R_1, R_2 > R_0)$ is given by,

$$P(R_1, R_2 > R_0) = P(R_2 > R_0/R_1 > R_0).P(R_1 > R_0) \quad (1)$$

where R_1 and R_2 are the rainfall rate in the stations 1 and 2, respectively; $P(R_1 > R_0)$ is the probability that R_0 be exceeded in the station 1, which corresponds to the cumulative statistical distribution of rain in station 1; and $P(R_2 > R_0/R_1 > R_0).P(R_1 > R_0)$ is the probability that R_0 be exceeded in station 2, once $R_1 > R_0$.

On the other hand, the conditional probability $P(R_2 > R_0/R_1 > R_0).P(R_1 > R_0)$ is related to the union of two events without intersection: the probability of the same rain cell is over station 1 and station 2 or the probability of one rain cell over station 1 and the other over station 2.

These probabilities are related to the following functions: i) $p(D_i, R_0)$ – joint probability density function of D_i and R_0 ; ii) $p(D_i)$ – probability density function corresponding to a rain cell with a diameter D_i in a position covering the two stations.

Based on the knowledge of the horizontal distribution of rain in the area under study, expression (1) can be written as,

$$P(R_1, R_2 > R_0) = \left\{ \sum_{D_i=L}^{\infty} p(D_i, R_0).p(D_i) + \left[\left(\sum_{D_i=0}^L p(D_i, R_0) + \sum_{D_i=L}^{\infty} [p(D_i, R_0).(1-p(D_i))] \right)^2 P_R(R_0) \right] \right\} P_R(R_0) \quad (2)$$

where $P_R(R_0) = P(R_1 > R_0)$.

Two limit conditions can be used to validate expression (2). First, if L tends to zero, it is easy to show that $P(R_1, R_2 > R_0)$ tends to $P(R_1 > R_0)$. On the other hand, if L tends to ∞ , $P(R_1, R_2 > R_0)$ tends to $[P(R_1 > R_0)]^2$.

3. Z–R relation

Radar reflectivity factor (Z) in mm^6/m^3 is related to rainfall rate (R) in mm/h through an empirical relation of type $Z = aR^b$. Based on rain gauge [7.36°S ; 72.80°W] and radar measurements [7.36°S ; 71.56°W] carried out in Cruzeiro do Sul, it was found the following values for the parameters a and b ,

$$a = 207 ; b = 1.53$$

The relation $Z = 207 R^{1.53}$ will be used throughout in this paper. Details about the derivation of “ a ” and “ b ” are being published elsewhere [1]. Figure 2 shows a comparison between the cumulative distribution of precipitation rate from rain gauge measurements and the corresponding distribution obtained by the conversion of radar data. An excellent agreement is observed for percentage of time up to 0.005 %. For smaller percentage the difference of values between the two procedures is probably due to rain gage saturation.

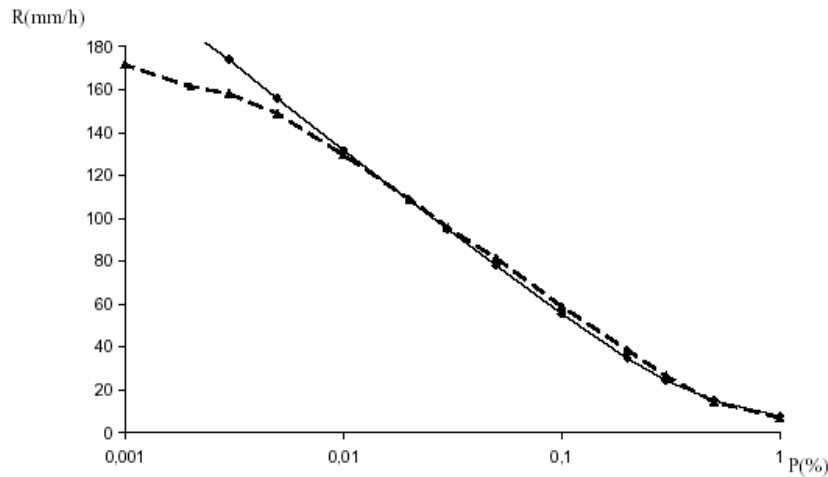


Figure 2. Comparison between cumulative distribution of precipitation rate from rain gauge (---) and radar data (—).

The analysis of the horizontal distribution of rain was based on the above relation between Z and R . The rain cell threshold was supposed to be 5 mm/h. First, it was derived the cumulative distribution of rain cell diameter for a given precipitation rate. This procedure allows to draw Figure 3, which shows the percentage of cells (vertical axis) with a diameter greater than the value indicated in the horizontal axis for $R \geq R_0$, being R_0 the reference for each curve.

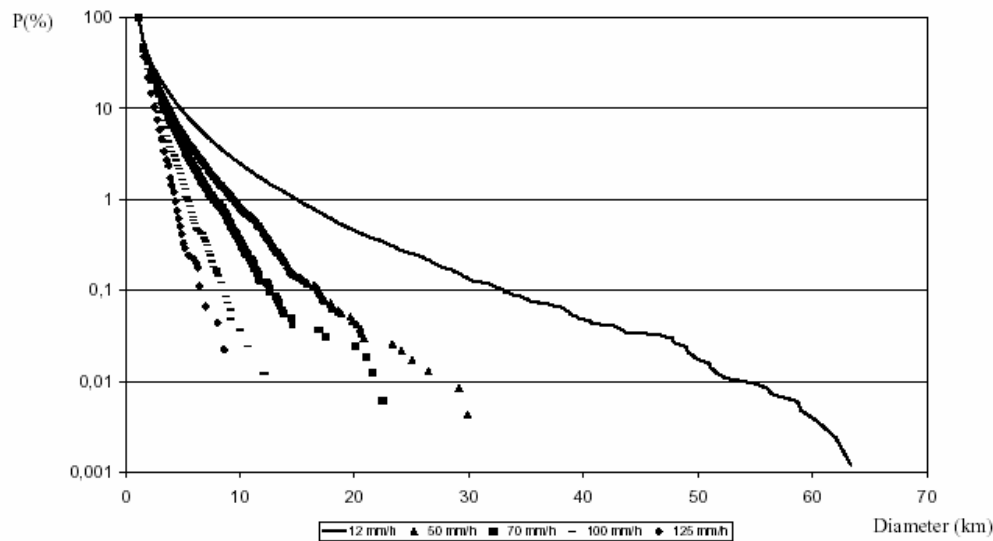


Figure 3. Rain cell diameter (km) exceeding a given precipitation rate (mm/h).

Finally, introducing the above results on equation (2), it was possible to draw the Figure 4, where it is shown the percentage of time that a precipitation rate greater than R_0 is exceeded in two stations separated by a given distance L . It is observed in this figure, for all curves, that a saturation occurs for distances between 10 and 15 km. This behavior is similar to site diversity gain curves empirically derived from rain attenuation measurements [2,3]. Once system parameters as, frequency, polarization and elevation angle, affect only the rain attenuation value, it is concluded that the saturation effect is due to rain correlation, which decreases as the separation between terminals increases.

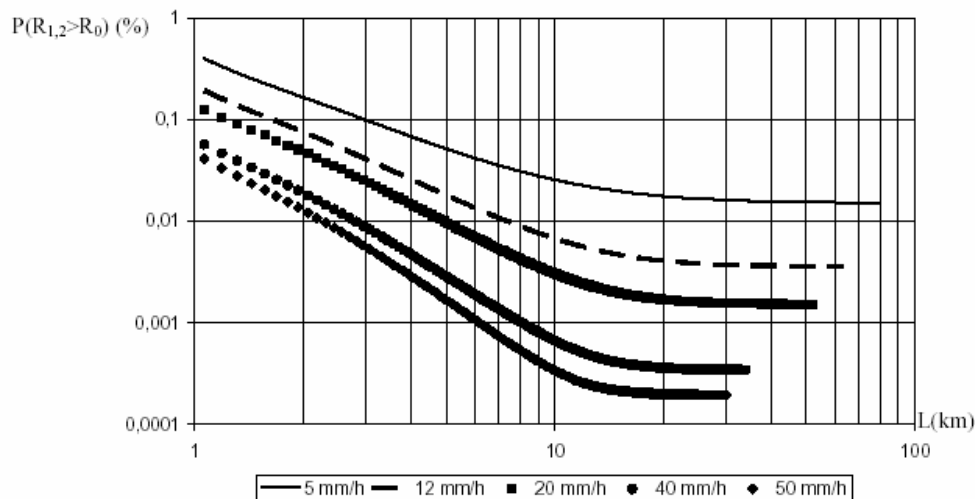


Figure 4. Percentage of time (%) that cumulative rain exceeding R_0 (mm/h) at two earth stations in a site diversity configuration, with distance L (km).

Based on this result, it is possible to settle an expression relating the rainfall rate (R_A) exceeded in station 1 for a given time percentage (p) with the rainfall rate (R_B) exceeded in both stations for the same time percentage. As a preliminary proposal, it can be written,

$$R_A = R_B (12.5 p + 1.1) L^{(0.42 \log(p) + 1.9)} \quad (3)$$

R_A and R_B in mm/h, L in km and p in %.

This preliminary expression will be checked with data from other radar stations.

4. Concluding remarks

Having as reference measurements from radar and rain gauges carried out in Cruzeiro do Sul [7.36°S ; 72.46°W], this paper has investigated the problem of spatial variability of rain in the Amazon region. A mathematical model relating the statistical distribution of rainfall rate between two points separated by a given distance was developed. The behavior of this model is similar to site diversity gain curves empirically derived from rain attenuation measurements. Based on this result, it is being proposed a practical formula relating the point rainfall rate with the rainfall rate simultaneously exceeded in two stations, in a site diversity configuration, for the same percentage of time.

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An Exceptional Rainfall Event in the Amazon Region

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Abstract

This paper describes an exceptional rainfall recorded at the radar station of Tefé [3.22°S ; 64.42°W] in the Amazon region. The rain cell dimension and rain height measured during this event show a large deviation from typical horizontal and vertical distributions. The effect on both terrestrial and slant paths is commented.

1. Introduction

The equatorial region is characterized by intense precipitation. Rainfall intensity levels around or above 100 mm/h are usually exceeded for more than 0.01% in the yearly cumulative distribution. However, this limit is, sometimes, surpassed with a very large margin and contributes for a non-expected behaviour in the small percentage of time range of the statistical distribution of rain. This paper describes an exceptional event where the horizontal distribution of rain covers a large area with very high intensity. The effect of this type of event on both terrestrial and slant paths are commented. Experimental data presented here are from a research program which is being developed under the responsibility of the Military Institute of Engineering (IME) with financial support from the National Scientific and Technological Development Council (CNPq). Figure 1 shows the rain gauge and radar networks used in this program, which are co-located and covering all the Amazon region.



Figure 1. Amazon region – rain gauge and radar stations.

2. Horizontal distribution of rain

Figure 2 shows a convective rain cell structure measured at the radar station of Cruzeiro do Sul [7.36°S ; 72.46°W]. Once the dimension of each small square in this figure is 1x1km, the mean cell diameter for $R = 100$ mm/h is around 1.0 to 1.5 km. Similar results have been observed throughout the network shown in Figure 1, and can be considered as typical for low latitude areas. The conversion from radar reflectivity (Z) to rainfall rate (R) was based on the following equation [1],

$$Z(\text{mm}^6/\text{m}^3) = 207[R(\text{mm/h})]^{1.53} \quad (1)$$

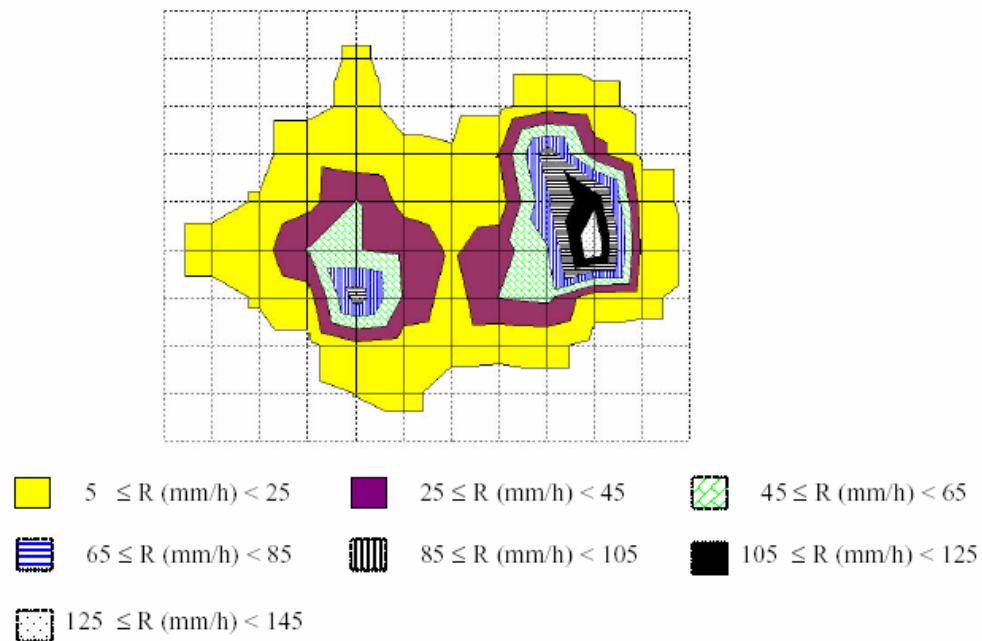


Figure 2. A typical example of the horizontal variability of rain.

Eventually, however, much higher rainfall rates, covering several km^2 , can be recorded in the Amazon region. An example is given in Figure 3, which was registered on 10/18/2004 by the meteorological radar located in Tefê [3.22°S ; 64.42°W]. In this example, the rain is distributed over a very broad area with an intensity higher than 200 mm/h. On the other hand, it should be pointed out that the measure of such precipitation rates with a rain gauge is quite difficult due to instrument saturation.

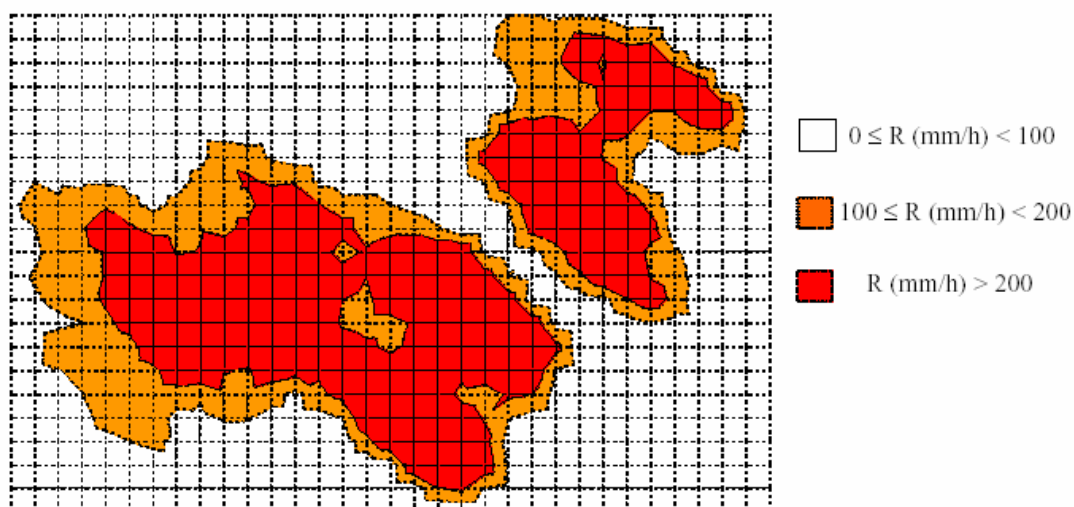


Figure 3. Horizontal distribution of an exceptional rainfall rate event.

3. Vertical distribution of rain

Along the years, the 0° isotherm h_0 has been used as reference for modeling the rain height. In general, experimental data available indicate that this reference has a sound basis. For instance, the profiles of radar reflectivity show in Figure 4 are in accordance with it. Sometimes, however, radar data point out the existence of heavy rain well above the 0° isotherm. An example is given in Figure 5 with data from the radar station of Tefé recorded also in the event of 10/18/2005. This figure shows vertical profiles of Z measured at 8 points separated by 250m in a range of 2km. Considering that h_0 is around 4.5km, the rain rate in this example is constant up to 9km.

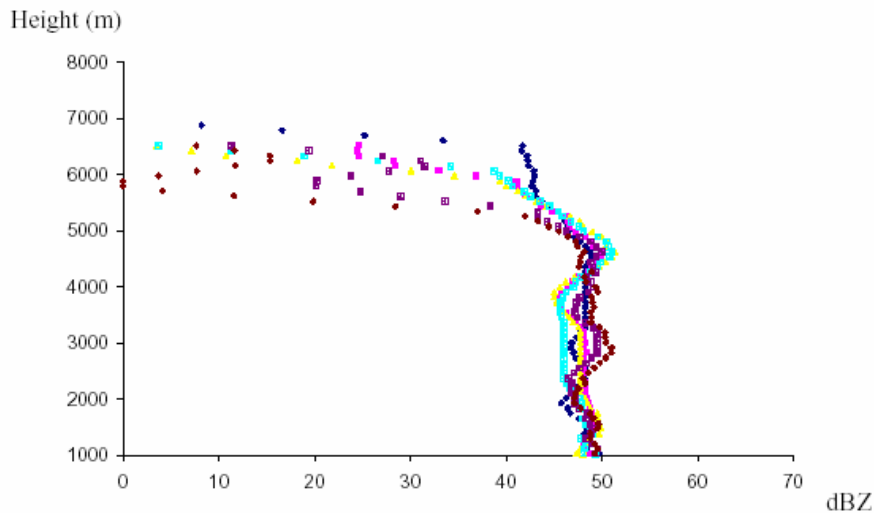


Figure 4. A typical example of the vertical distribution of rain.

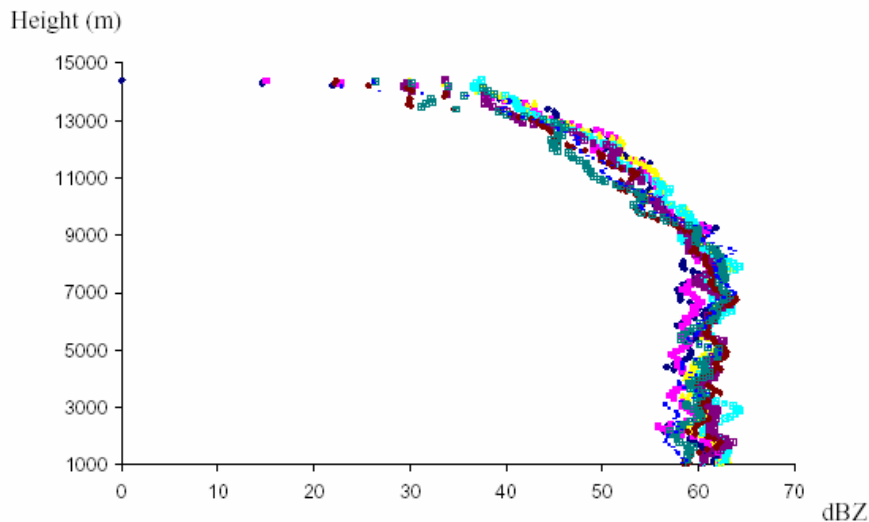


Figure 5. Vertical distribution of an exceptional rainfall event.

4. Effect on terrestrial and slant paths

Referring to terrestrial paths, when such type of event occurs, a very high attenuation level is noticed. Considering that the rain cell of Figure 3 covers a distance of more than 15km with a rain intensity higher than 200 mm/h, a typical radio link (5 ~ 10km) could be totally immersed on it.

On the other hand, a study by Timóteo da Costa and Assis [2] has shown large errors when comparing rain height corresponding to -2°C isotherm height, as adopted in the IUT-R model [3], with the effective rain height derived from experimental data measured in low latitude areas. Possibly, this disagreement is due to the vertical distribution of rain commented above.

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Análise da Variação Espacial da Taxa de Precipitação com Aplicação em Diversidade de Sítio para Enlaces Terra-Satélite

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Resumo—Este trabalho analisa a variação espacial da taxa de precipitação na região Amazônica. Foram utilizadas medidas de um radar meteorológico localizado Cruzeiro do Sul, AC, e desenvolvido um modelo matemático relacionando a distribuição estatística da chuva entre dois pontos separados por uma determinada distância. Adicionalmente, no estabelecimento da relação entre refletividade do radar e taxa de precipitação, dados provenientes de um pluviógrafo instalado nas vizinhanças do radar foram levados em conta. Os resultados finais do trabalho mostraram-se compatíveis com a forma de variação observada nas curvas de ganho da diversidade de sítio obtidas a partir de modelos empíricos com base na atenuação do sinal.

Palavras-Chave— taxa de precipitação, diversidade de sítio, região Amazônica, radar meteorológico, radiometeorologia,

Abstract—This paper deals with the spatial variability of rain rate in the Amazon region. Measurements from a meteorological radar located in Cruzeiro do Sul, AC, were taken into account and a mathematical model for relating the statistical distribution of precipitation rate between two points separated by a given distance was developed. Additionally, data from a pluviograph installed near the radar was also used in the derivation of a relationship between radar reflectivity and precipitation rate. Final results were promising enough. The behaviour of the mathematical model has showed excellent agreement with site diversity gain curves from empirical models based on rain attenuation.

Keywords— precipitation rate, site diversity, Amazon region, meteorological radar, radiometeorology.

I. INTRODUÇÃO

A diversidade de sítio representa a solução técnica mais eficiente para compensar o efeito da atenuação por chuva em enlaces terra-satélite que operam em frequências acima de

10 GHz. É claro que esta solução implica em maior custo na implementação do sistema. Entretanto, em locais onde o emprego de comunicações via satélite seja a melhor alternativa, a relação custo-benefício poderá justificar o investimento realizado. A Amazônia é um exemplo a ser citado, uma vez que as distâncias envolvidas e a barreira natural da floresta indicam o satélite como um caminho natural para resolver os problemas de comunicações. Por outro lado, o uso da banda Ku (12 – 18 GHz) nesta região poderá ser inviabilizado se não houver uma utilização criteriosa de técnicas que possibilitem a operação dentro de padrões aceitáveis de disponibilidade.

Com base em medidas de radar realizadas na localidade de Cruzeiro do Sul, divisa do Acre com Amazonas, este trabalho analisa o problema da variação espacial da chuva na região equatorial. Trata-se de um passo importante no sentido de desenvolver, a partir de uma sólida base física, um modelo que permita avaliar o desempenho da diversidade de sítio. A segunda parte desta pesquisa, em andamento, será estabelecer um procedimento de cálculo da atenuação por chuva aplicável nas condições de interesse na Amazônia, quais sejam, altas taxas de precipitação e ângulos de elevação acima de 60°. Cumpre ressaltar que, em tais condições, o modelo [1] adotado atualmente pelo Setor de Radiocomunicações da União Internacional de Telecomunicações (UIT-R) apresenta um comportamento aparentemente inadequado [2]. Esta questão está sendo investigada e os resultados obtidos serão publicados oportunamente.

Os dados experimentais aqui apresentados são provenientes de um amplo programa de pesquisa em radiometeorologia que está sendo realizado na região amazônica, sob a responsabilidade do Instituto Militar de Engenharia (IME), com apoio do Conselho Nacional de Desenvolvimento Científico e Tecnológico (CNPq), da Pontifícia Universidade Católica do Rio de Janeiro (PUC-Rio) e do Sistema de Proteção da Amazônia (SIPAM). Este programa tem como base de medidas a rede de radares e pluviógrafos mostrada na Figura 1.

II. FORMULAÇÃO MATEMÁTICA DO PROBLEMA

A questão fundamental que se apresenta é determinar a separação que deve existir entre as duas estações terrenas de uma configuração em diversidade para que o efeito da chuva seja minimizado e permita que a disponibilidade seja adequada

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para o serviço que está sendo prestado. Para isto torna-se necessário avaliar a probabilidade de que uma determinada taxa de precipitação (R_0) seja excedida, simultaneamente, nas duas estações distanciadas por uma distância L_0 .

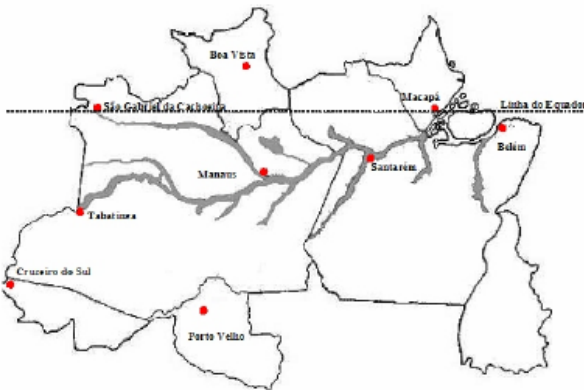


Fig. 1. Rede de Pluviômetros e Radares Meteorológicos.

Esta probabilidade pode ser determinada através de uma probabilidade condicional:

$$P(R_1 > R_0 \text{ e } R_2 > R_0) = P(R_2 > R_0 / R_1 > R_0) \cdot P(R_1 > R_0) \quad (1)$$

onde R_1 e R_2 são as taxas de precipitação na estação 1 e 2, respectivamente.

A probabilidade da chuva ser excedida na estação 1, $P(R_1 > R_0)$, será dada pela função distribuição cumulativa da chuva na região, $F_R(R_0)$.

A probabilidade da chuva da estação 2 exceder a chuva especificada, dado que a chuva na estação 1 está excedida, $P(R_2 > R_0 / R_1 > R_0)$, pode ser determinada pela união de dois eventos disjuntos, isso é, a probabilidade das estações 1 e 2 serem atingidas pela mesma célula de chuva (p_{mc}) ou a probabilidade de serem atingidas por células distintas (p_{cd}).

Nos cálculos destas probabilidades, leva-se em conta a função densidade de probabilidade do tamanho horizontal da célula de chuva para uma determinada taxa de chuva excedida, $f_D(D, R_0)$. Considerando a seção horizontal de uma célula de chuva como um círculo, este tamanho será definido pelo diâmetro da célula de chuva, D .

A probabilidade da estação 2 ser afetada pela mesma célula de chuva, excedendo R_0 , que está atingindo a estação 1, será dada pelo produto entre probabilidade do diâmetro da célula ser D_i , $f_D(D_i, R_0)$, e a probabilidade do centro desta célula esteja em pontos específicos, isso é, pontos que permitam atingir ambas as estações simultaneamente, $p_1(D_i)$. O valor de $p_1(D_i)$ será dado por uma relação de áreas.

A probabilidade da estação 2 ser afetada por uma célula de chuva, excedendo R_0 , distinta da que atinge a estação 1, será dada pelo produto entre a probabilidade de existir esta nova célula de chuva, $F_R(R_0)$, e a probabilidades de ambas as

células apresentem diâmetros D_{i1} e D_{i2} e estejam em centros específicos de forma a não atingirem, cada uma, as duas estações.

Assim, a equação 1 será determinada por:

$$P(R_1 > R_0 \text{ e } R_2 > R_0) = \left\{ \sum_{D_i=L_0}^{\infty} f_D(D_i, R_0) p_1(D_i) + \right. \quad (2)$$

$$\left. + \left[\left(\sum_{D_i=0}^{L_0} f_D(D_i, R_0) + \sum_{D_i=L_0}^{\infty} [f_D(D_i, R_0) \cdot (1 - p_1(D_i))] \right) F_R(R_0) \right] \right\} F_R(R_0)$$

Pode ser verificado que o valor da probabilidade dada pela equação (2) tende para $F_R(R_0)$ quando a distância entre as duas estações, L_0 , tende para zero. Quando esta distância tende para o infinito, isso é, para valores maiores que o diâmetro máximo de chuva encontrado na região, o valor da probabilidade dada pela equação (2) tende para o quadrado de $F_R(R_0)$.

III. LEVANTAMENTO DE DADOS E ANÁLISE DAS MEDIDAS

O primeiro passo para o cálculo das probabilidades associadas à estrutura horizontal da chuva, foi o de se estabelecer uma relação entre a refletividade (Z) do radar e a taxa de precipitação (R). Para isto, foram utilizadas medidas realizadas no período de agosto a dezembro de 2004, totalizando 6.908 imagens do tipo CAPPI ("Constant Altitude Plan Polar Indicator"), provenientes do radar de Cruzeiro do Sul, obtidas a uma altura em relação ao solo de 2 km e com uma distância máxima ao radar de 80 km. Este valor de altura foi utilizado por ser um valor intermediário entre o solo e a altura da célula de chuva, considerando também que a variação da refletividade (Z) com a altura no percurso da chuva não é relevante [3]. Ao final de uma varredura CAPPI dispõe-se de um valor de Z em cada quadricula de dimensão 1 km x 1 km. De posse de tais medidas, utilizando como referência a relação de Marshall e Palmer [4], ou seja,

$$Z(\text{mm}^6/\text{m}^3) = 200[R(\text{mm/h})]^{1.6} \quad (3)$$

e tendo como elemento de comparação à distribuição cumulativa da taxa de precipitação medida no pluviômetro localizado nas vizinhanças do radar, chegou-se a

$$Z(\text{mm}^6/\text{m}^3) = 207[R(\text{mm/h})]^{1.53} \quad (4)$$

Os detalhes da obtenção de (4) estão descritos em outra publicação [5]. Observa-se que a relação é muito próxima de Marshall e Palmer. Isto era esperado, uma vez que, para o objetivo pretendido neste trabalho, considerou-se o limiar da taxa de precipitação em 5 mm/h, abrangendo, desta forma, tanto chuva estratiforme, como chuva convectiva. A figura 2 apresenta a distribuição cumulativa da taxa de precipitação medida pelo pluviômetro no período de janeiro a dezembro de 2004 e aquela obtida a partir da equação (4).

A seguir, associando a cada quadricula o valor de R correspondente, considerou-se como chuva toda quadricula com taxa de precipitação igual ou superior a 5 mm/h. Definiram-se as células de chuva excedendo R_0 , como o conjunto de quadriculas, com $R > R_0$, que apresentasse ligação horizontal ou vertical entre seus lados. A área de cada célula foi determinada pelo número de quadriculas existentes.

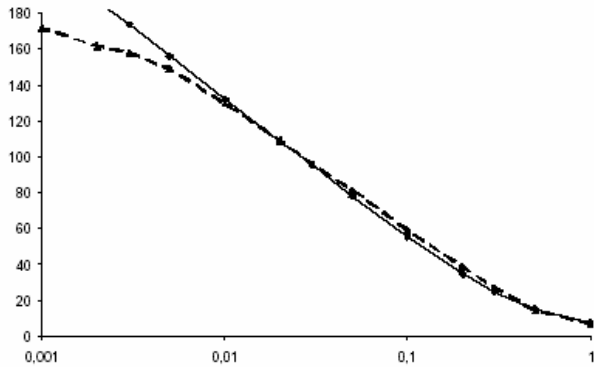


Fig. 2. Função distribuição acumulada da chuva em Cruzeiro do Sul (F_R) medida pelo pluviômetro (---) e através da equação 4 (—). Eixo horizontal marca a probabilidade (%) e o eixo vertical o valor da taxa de precipitação cumulativa excedida (mm/h).

Fazendo equivalência da área irregular assim obtida com uma circunferência, chegou-se a uma distribuição dos valores de diâmetro da célula para cada taxa de precipitação excedida. Assim foi possível traçar a Figura 3 que apresenta a percentagem do número de células (eixo vertical) com diâmetro superior ao valor indicado no eixo horizontal para $R \geq R_0$, sendo R_0 o limiar tomado como referência em cada curva.

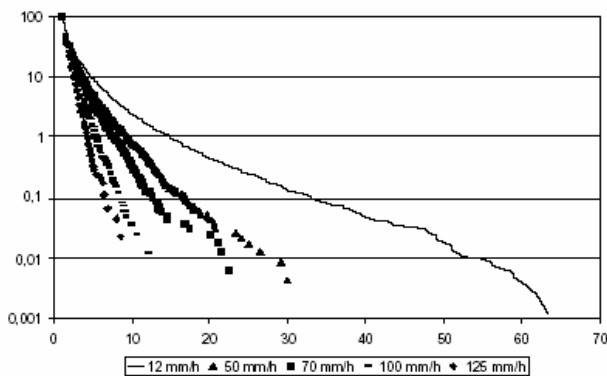


Fig. 3. Percentagem do número de células (%), eixo vertical, com diâmetro (km) superior ao valor indicado no eixo horizontal, para diversos valores de $R \geq R_0$, sendo R_0 o limiar tomado como referência em cada curva.

Por fim, baseado na equação (2), obtém-se então a Figura 4 que mostra a percentagem de tempo que uma taxa de precipitação acima de R_0 é excedida em duas estações separadas por uma determinada distância, L_0 .

Verifica-se nesta figura que existe uma saturação das curvas entre 10 e 15 km. Este resultado é similar ao que se observa nas curvas de ganho de diversidade obtidos a partir de modelos empíricos que tem por base a atenuação do sinal [1,6]. Considerando que os parâmetros de sistema, tais como, frequência, polarização e ângulo de elevação afetam unicamente o valor da atenuação, conclui-se que o efeito de saturação é função apenas da decorrelação da chuva à medida que aumenta a distância de separação entre os terminais.

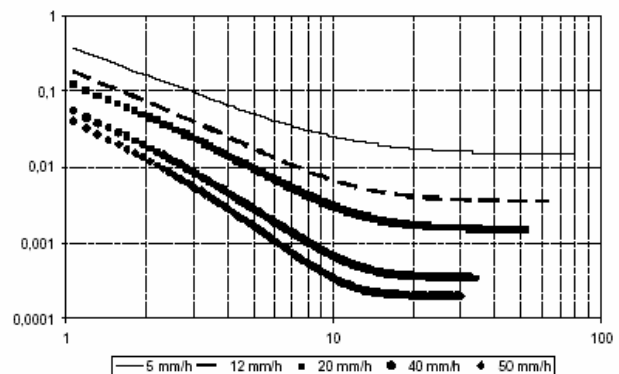


Fig. 4. Percentagem de tempo (%), eixo vertical, que uma taxa de precipitação acima de R_0 (mm/h) é excedida em duas estações separadas por uma determinada distância L_0 (km), eixo horizontal.

Em vista deste resultado, procurou-se definir um relacionamento entre a taxa de chuva excedida em uma estação isoladamente (R_A) em uma dada percentagem de tempo (p) e o valor da taxa excedida simultaneamente em duas estações (R_B), separadas por uma distância da na mesma percentagem de tempo. O melhor ajuste correspondeu a,

$$R_A = R_B \cdot (12,5 p + 1,1) D^{(0,42 \log(p) + 1,9)} \quad (5)$$

sendo R_A e R_B em mm/h, D em km e p em %.

A expressão (5) é considerada ainda preliminar, pois se pretende investigar dados de outras estações na mesma linha do que foi feito neste trabalho. Pode-se, no entanto, adiantar que se espera um comportamento similar, pois as características da chuva na região em estudo não apresentam variações significativas.

IV. CONCLUSÕES

Tendo por referência medidas (radar e pluviômetro) realizadas em Cruzeiro do Sul, AC, este trabalho analisou o problema da variação espacial da taxa de precipitação na região Amazônica. A formulação matemática desenvolvida para modelar esta variação mostrou-se plenamente

satisfatória. A saturação observada na figura 4 do trabalho é similar ao comportamento observado nas curvas empíricas de ganho de diversidade, demonstrando que o modelo definido pela equação (2) é adequado para análise da variação espacial da taxa de precipitação. Com base neste resultado foi proposta uma fórmula prática a ser utilizada no dimensionamento da separação entre terminais de um sistema de diversidade de sítio.

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RAIN ATTENUATION RESEARCH IN BRAZIL

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ABSTRACT

This paper describes some preliminary results and the activities in progress of a research program on rain attenuation in the Amazon region. This program is part of a broad project, named CT-Amazônia, covering different aspects of science and technology, which is being developed under the responsibility of the Military Institute of Engineering (IME) with financial support from the National Scientific and Technological Development Council (CNPq). The following topics are being investigated in this program: a) Precipitation rate behaviour based on rain gauge and radar measurements and b) Rain attenuation prediction modelling for terrestrial and slant paths.

INTRODUCTION

The research on rain attenuation in Brazil has started around 1970 [1]. Reference to the work prior to 1996 may be found in [2] and [3]. This paper presents some preliminary results related to rainfall characteristics, as well as, the research program now in progress. This research is the theme of the doctoral thesis of the first author and includes two main topics: a) Precipitation rate distribution and radar measurements in the Amazon region and b) Rain attenuation prediction models for terrestrial and slant propagation paths. Referring to rainfall rate studies, a very broad rain gauge network was implemented in the Amazon region. Figure 1 shows the locations where the gauges were placed. On other hand, considering the well-known relation between radar reflectivity and rain rate, measurements from a radar network is being used as a complement for these studies. This radar network belongs to SIPAM (Amazon system protection) and covers the same locations as the rain gauge network. Radar data will be also used in the definition of rain cell shape and dimension. With this procedure, it is hoped to derive an attenuation model closely fitted to physical reality of rain structure.

KÖPPEN CLIMATE CLASSIFICATION

According to preliminary results of this research program, the combined use of precipitation and radar data, Köppen climate classification [4] and the mathematical model adopted by the ITU-R Recommendation P.837-4 [5] appears to be an useful tool for the prediction of rainfall rate cumulative probability distribution. This approach can also be used to improve contours of homogeneous climatic areas, as well as to better characterize the transition between different climates.

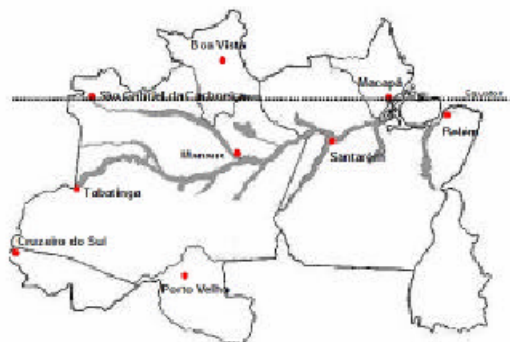


Fig. 1 Amazon region – rain gauge and radar networks.

Köppen classification was adopted because its structure depends on temperature, precipitation and vegetation, factors that can be related to the statistical distribution of rain in a given area. In the Amazon region only Köppen climate A (tropical rainy) can be identified with the following subtypes:

- Rainy Equatorial (Af) - with a large annual rainfall (over 2000 mm) and practically no dry season;
- Monsoon Tropical (Am) – the annual rainfall is equal to or larger than Af, but there is a short dry season (one to three months);
- Wet-and-dry tropical (Aw) – where the rain and dry seasons are well defined.

PRECIPITATION RATE DISTRIBUTION

When there is no information about the local cumulative distribution of rain, an alternative is to use a mathematical model based on meteorological parameters available in the region under study. In this paper it was adopted the Salonen-Baptista model, which is the basis of the ITU-R Recommendation P.837-4 [5]. This model depends on the following parameters:

- ❖ Average annual convective rainfall amount, M_C (mm);
- ❖ Average annual stratiform rainfall amount, M_S (mm);
- ❖ Probability of rainy 6h periods, Pr_6 .

which can be obtained from the data available in the Brazilian Meteorological Institute.

Taking this mathematical model together with precipitation and radar data, a detailed analysis of the cumulative distribution of rain in the Amazon region is being carried out. Figure 2 shows a comparison between the cumulative distribution measured at Cruzeiro do Sul, from January 2004 to December 2004, with the corresponding predicted values from the Salonen-Baptista model adjusted with meteorological local data. An acceptable agreement between the two curves is observed.

Z – R RELATION

Radar reflectivity factor (Z) in mm^6/mm^3 is related to rainfall rate (R) in mm/h through an empirical relation of the type $Z = aR^b$. Based on rain gauge [7.36°S; 72.80°W] and radar measurements [7.36°S ; 71.56°W] carried out at Cruzeiro do Sul, it was found the following values for these parameters: $a = 207$ and $b = 1.53$.

The relation $Z = 207 R^{1.53}$ will be used throughout in this paper. Details about the derivation of “a” and “b” are being published elsewhere [6]. Figure 3 shows a comparison between the cumulative distribution of precipitation rate from rain gauge measurements and the corresponding distribution obtained by the conversion of radar data. An excellent agreement is observed for percentage of time up to 0.005 %. For smaller percentages of time the difference of values between the two procedures is probably due to rain gauge saturation.

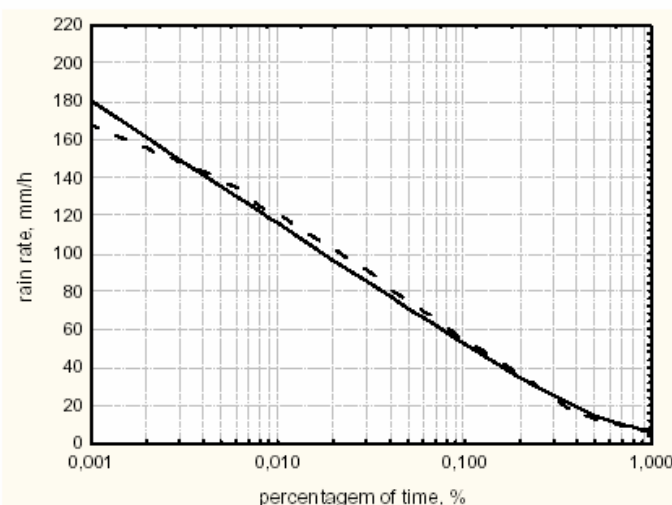


Fig. 2 Cumulative distribution of rainfall rate in Cruzeiro do Sul (7.36°S ; 72.80°W) - Comparison between rainfall rate measurements (—) and Salonen-Baptista prediction model (- - -).

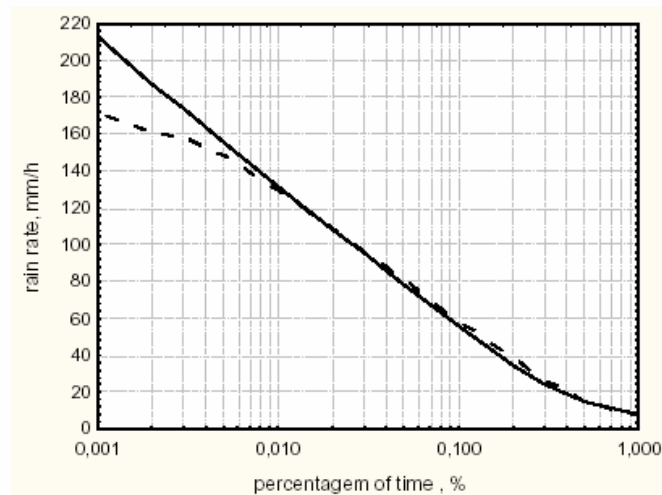


Fig. 3 - Cumulative distribution of rainfall rate in Cruzeiro do Sul - Comparison between rainfall rate measurements from rain gauge (---) and radar data (—)

PREDICTION OF RAIN ATTENUATION

An example of a rain cell measured at the radar station of Cruzeiro do Sul is given in Figure 4. Taking into account the non-uniform aspect of the rain cell shown in this figure, it is clear the difficulty to model it by a simple geometrical shape. Most of prediction models use the concept of path length factor (or distance factor) for solving the problem of non-uniformity of rain along the propagation path. However, considering that is not feasible to derive a rigorous mathematical expression for this factor, an alternative is the use of experimental data for fitting an empirical solution based on a given rain cell model. Taking as reference a truncated exponential rain cell, promising results were obtained by Timóteo da Costa and Assis [7]. An important reason for choosing this model was because, through a combination of cylindrical and exponential shapes, it is avoided the restriction associated to a measured path length factor greater than one, common to other predicting methods.

Observing Figure 4, once the rotational symmetry assumed above is not in accordance with the shape of the rain cell, a limitation to this model could be raised. However, a detailed analysis of the cell core, where precipitation is more intense, indicates that this approximation seems to be acceptable. There is not a good agreement in the tail of the rain cell, where the precipitation is low and, consequently, less important for rain attenuation.

On the other hand, weather radar observations [8] show that, on average, the rain intensity does not vary from the surface of the Earth up to the 0°C isotherm (h_0). Based on this result, h_0 has been used along the years as reference for modelling the rain height. However, in the Amazon region (and probably in other low latitude locations) this point deserves a more detailed investigation, once radar data indicate the existence of heavy rain well above the 0°C isotherm. An example is given in Figure 6 with data from the radar station of Tefê [3.22°S ; 64.42°W]. This figure shows vertical profiles of Z measured at 8 points separated by 250 m in a range of 2 km. Considering that the isotherm of 0°C is around 4500m, the rain is approximately constant up to $2 h_0$, i. e. 9000m.

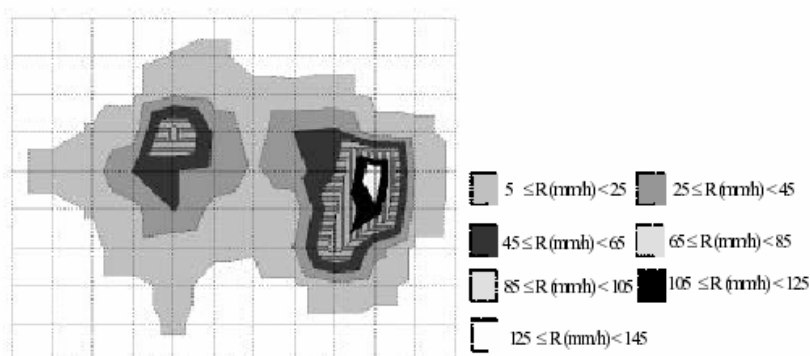


Fig. 4 – An example of the horizontal variability of rain

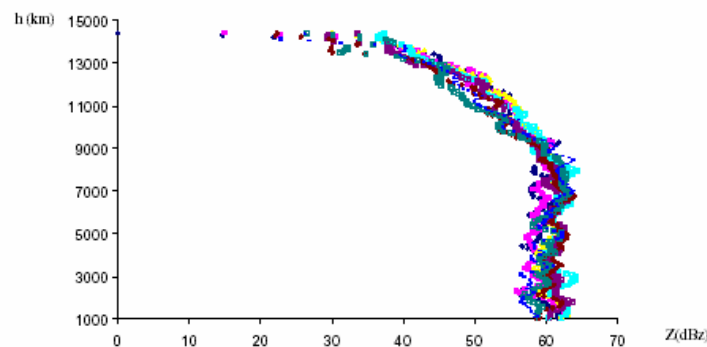


Fig. 5 An example of the vertical variability of rain.

The concept of equivalent rain height seems to be an adequate way to solve this question. However, as it was pointed out by Assis [9], this solution should be considered only after solving the problem of the horizontal distribution of rain. An study by Timóteo da Costa and Assis [10] has shown large errors when comparing rain height corresponding to -2°C isotherm height, as adopted in the ITU-R model [11] with the effective rain height derived from experimental data in latitudes between 30°N and 30°S .

CONCLUDING REMARKS

According to the preliminary results described in this paper, it is clear that the combined use of precipitation and radar data, Köppen climate classification and Salonen-Baptista mathematical model appears to be an useful tool for the prediction of rainfall rate cumulative distribution. This approach can also be used to improve contours of an homogeneous climatic area, as well as, to better characterize the transition between different climates. As expected for convective rain, the non-uniformity of precipitation in the horizontal plane constitutes an obstacle for modelling the cell. In spite of that, considering that the core of the rain cell can be treated as homogeneous, a model referenced by a truncated exponential rain cell is being investigated. Regarding vertical variability, once for convective precipitation radar signals can be detected well above the 0°C isotherm, the concept of effective rain height seems to be an adequate way to face the question. However, taking into account the totally empirical basis of this concept, it must be used only after the problem of horizontal distribution has been solved.

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